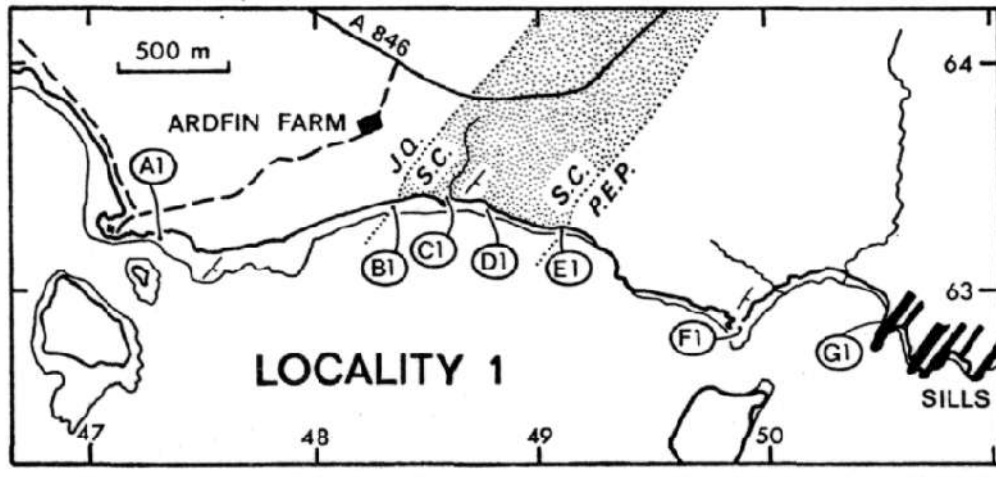


Excursion at Ardfin

This excursion is currently being updated following the construction of the golf course. Below is an extract from ANDERTON, R. (1977): *The Dalradian rocks of Jura. Scottish Journal of Geology*, **13**, pp. 135–142.

The shoreline traverse from A1 to G1 remains accessible. Park on the main road and walk down to A1, follow the coast eastwards and exit via the golf course roads near F1 back to the road.



Locality 1: Ardfin

A traverse from the Jura Quartzite through the Scarba Conglomerate into the Port Ellen Phyllites (Fig. 3).

Take the track from the main road down through Ardfin Farm towards the shore, having first obtained permission to continue at the farm. From (A1) east to the top of the Jura Quartzite there are some silts and laminated sands as well as the medium-grained to pebbly, cross-bedded sands which form the rather monotonous section from (A1) NW. along the Sound of Islay. A 4 m unit of light grey slate marks the base of the Scarba Conglomerate at (B1) and is followed eastwards by patchy exposures of deformed slates and cleaved quartzites. The slate darkens in colour and the proportion of sand decreases eastwards. At (C1), just east of a Tertiary dolerite dyke, thickly bedded sands appear. They differ from those in the Jura Quartzite in having thicker beds, more persistent bedding planes, occasional grading and small shale clasts. The next section of the shore (D1–E1) consists of fallen boulders from Tertiary dolerite dykes with a few exposures of graded sands in the cliff. At (E1) is the abrupt boundary with the Port Ellen Phyllites. The Phyllites consist of highly deformed green and black slates and phyllites with thin, sometimes calcareous, sandstones. The proportion of sand increases eastwards to (F1) where 15 m of cleaved, fine grained, green sandstone is exposed. Some beds have a massive base and a cleaved top suggesting that they are graded. A fine series of epidiorite sills form the headlands east of (G1). Returning west the relationship between lithology and the intensity of deformation can be considered.